

Revisiting of the greenhouse effect discovery and its analysis

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Outline

1. A short history of the greenhouse effect discovery
2. Analysis of the greenhouse effect using the Net Exchange Formalism

Emergence of the physics of climate

J. Fourier:

- *Mémoire sur les températures du globe terrestre et des espaces planétaires*, Mémoires de l'Académie des Sciences de l'Institut de France, 1824
- *General remarks on the Temperature of the Terrestrial Globe and the Planetary Spaces*; American Journal of Science, Vol. 32, N°1, 1837.



Joseph Fourier
(1768-1830)

- He consider the Earth like any other planet
- The energy balance equation drives the temperature of all the planets
- The major heat transfers are
 1. Solar radiation
 2. Infra-red radiation
 3. Diffusion with the interior of Earth
- The heat exchange with the interior of Earth has a negligible impact on the surface temperature

Emergence of the physics of climate

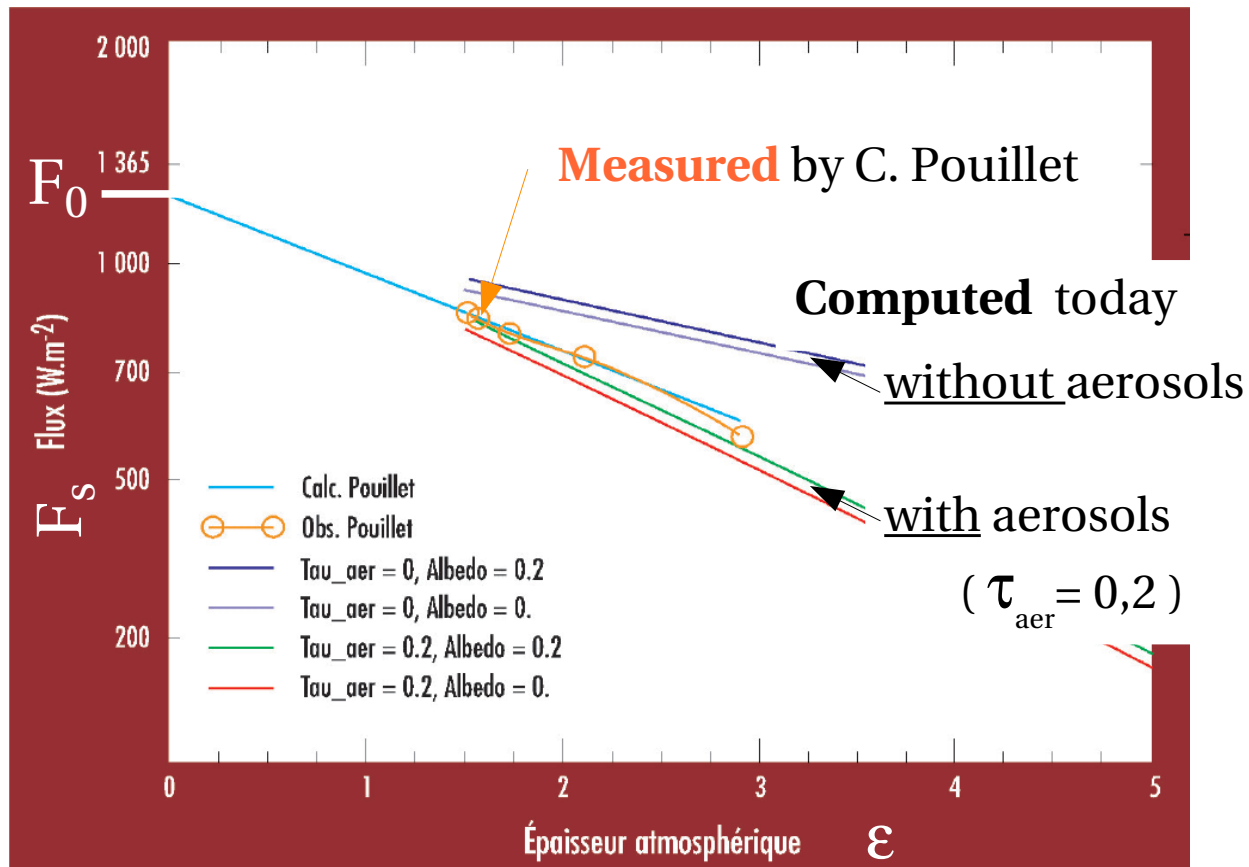
- He highlight the importance of solar and infrared radiation and makes an analogy between heat exchanges in the atmosphere and in a « hot box » (a greenhouse)
- He assume that the polar ward heat transport is small and deduced that, during the polar winter night, the polar temperature is equal to the temperature of the « *planetary space* »
- He envisage the importance of any change of the sun « *The least variation in the distance of that body[the sun] from the earth would occasion very considerable changes of temperature.* »
- He envisage that climate may change: « *The establishment and progress of human society, and the action of natural powers, may, in extensive regions, produce remarkable changes in the state of the surface, the distribution of waters, and the great movements of the air. Such effects, in the course of some centuries, must produce variations in the mean temperature for such places* ».

Estimate of the solar irradiance

Mémoire sur la chaleur solaire, sur les pouvoirs rayonnants et absorbants de l'air atmosphérique, et sur la température de l'espace, Pouillet, 1838

- Developed a pyrheliometer to measure the solar irradiance

Solar irradiance at surface F_s as a function of atmospheric air mass ϵ



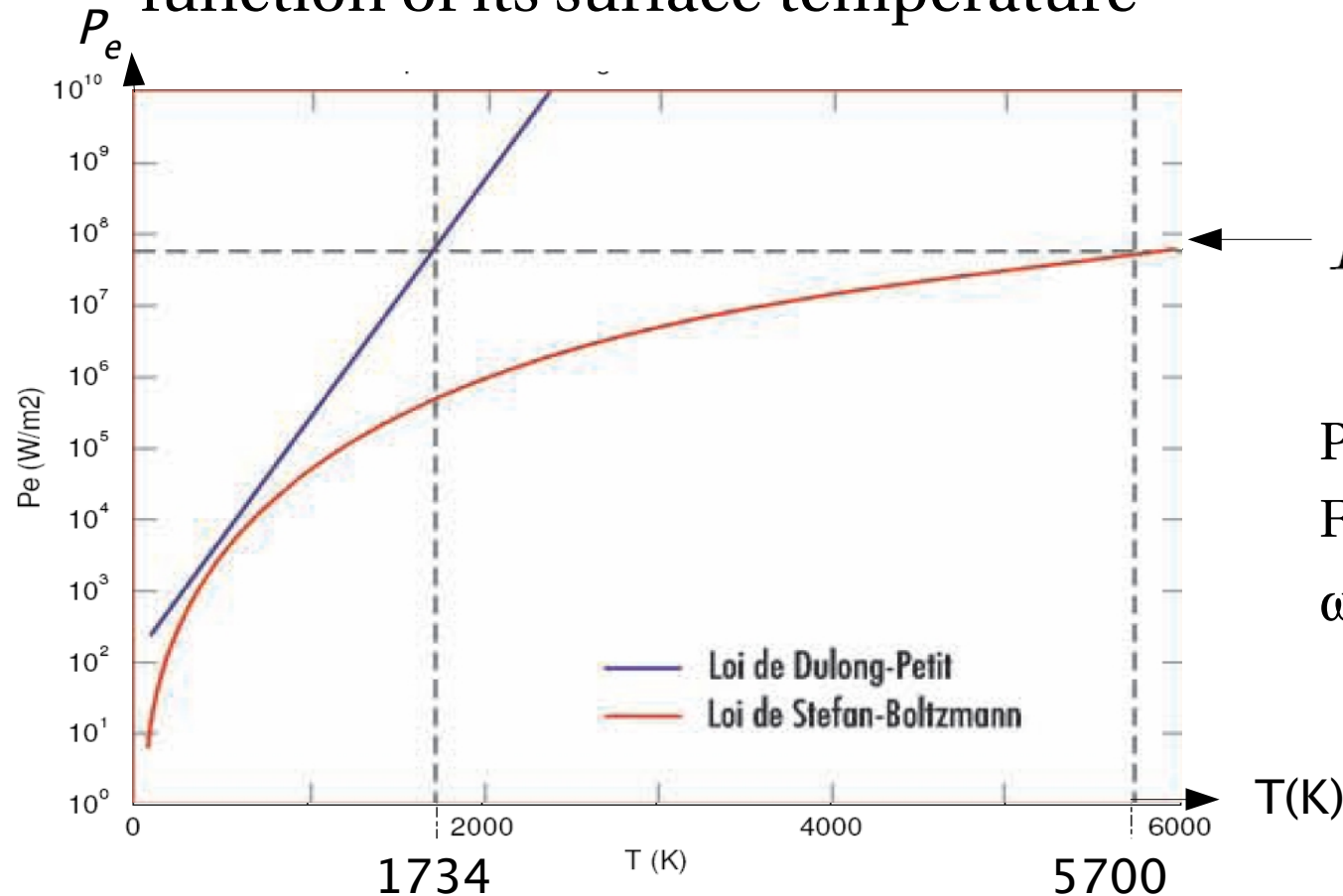
Claude Pouillet
(1790-1868)

C. Pouillet obtains:
 $F_0 = 1228 \text{ W.m}^{-2}$

Current estimate:
 $F_0 = 1361 \text{ W.m}^{-2}$

... and of the surface temperature of the sun

Flux emitted by the sun as a function of its surface temperature



Pouillet estimate

$$P_e = \frac{F_0}{\sin^2 \omega}$$

P_e : flux emitted by the sun

F_0 : solar constant

ω : angular radius of the sun

Same deduction, using the Stefan Boltzmann law

Deduced by C. Pouillet, using the Dulong-Petit law

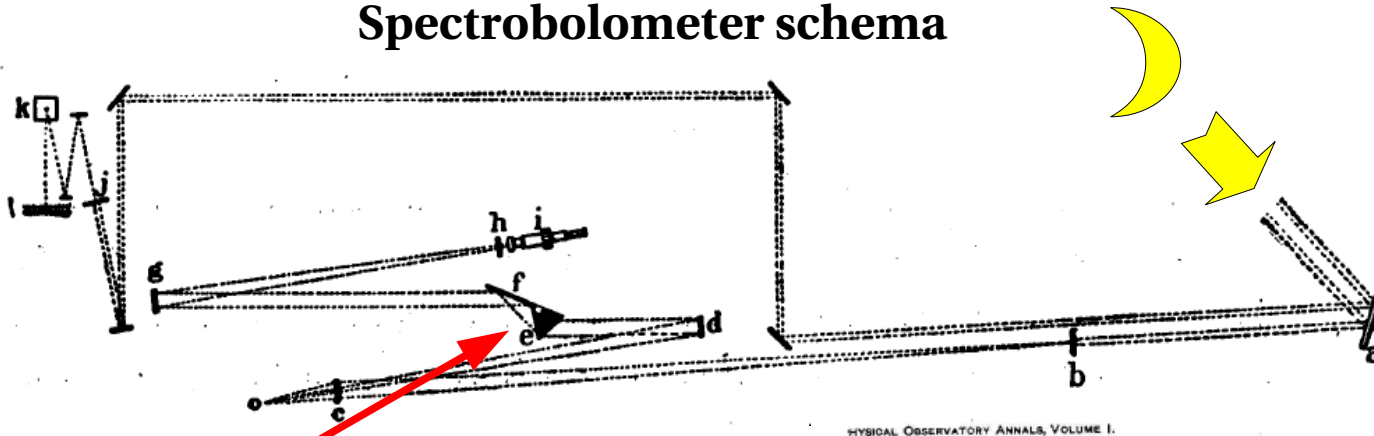
Measurement of the infrared spectra

- Developed the spectrobolometer
- Measured the solar spectra
- Measured the infrared radiation coming from the moon



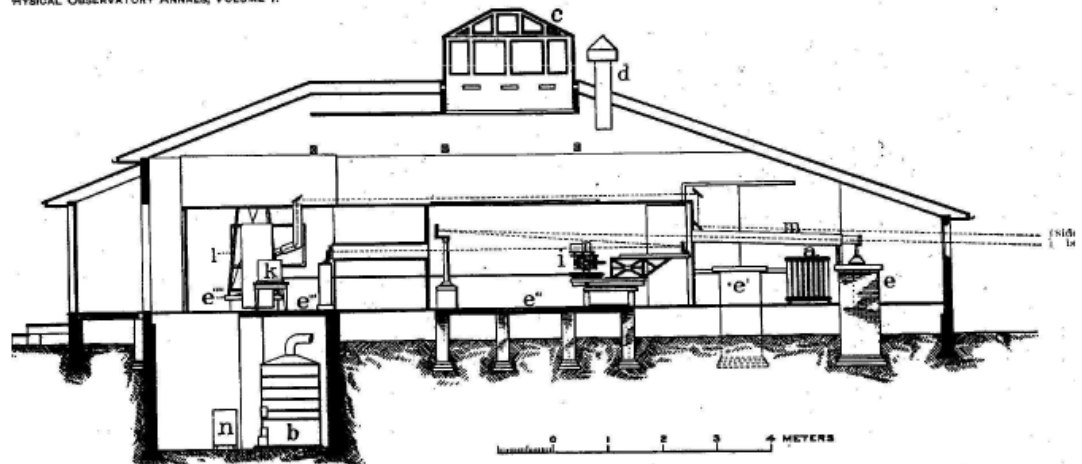
Samuel Langley
(1834-1906)

Spectrobolometer schema



Rock salt prism (NaCl)

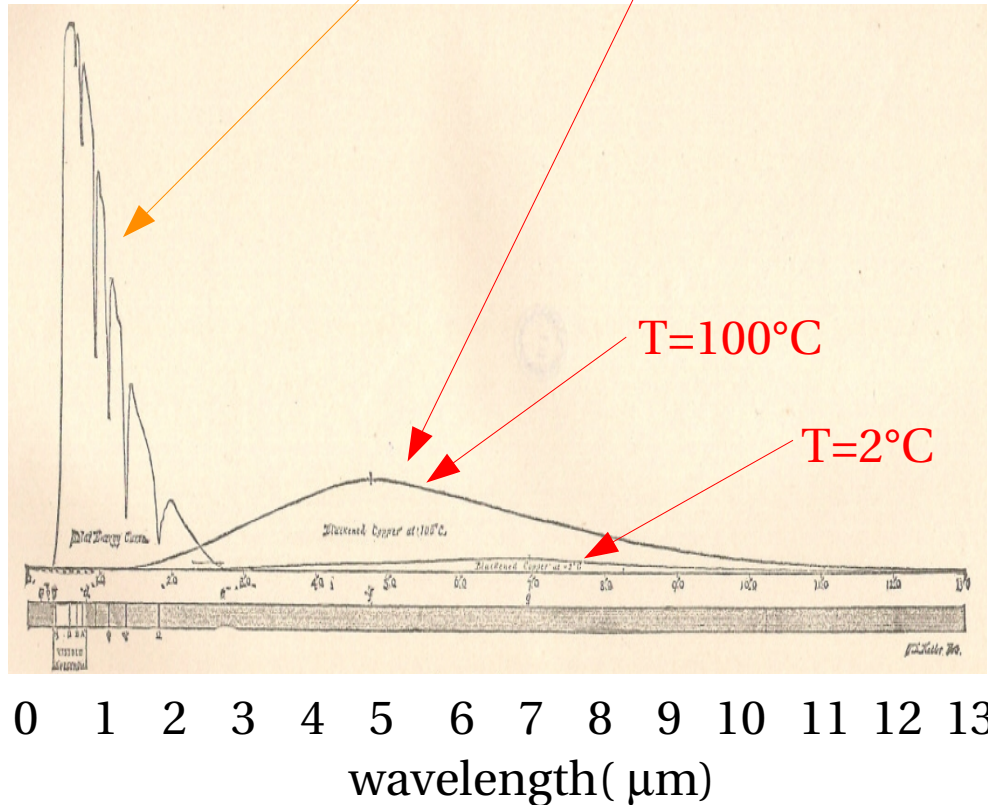
Going from deviation angles to wavelength require refractive index values, which were not known



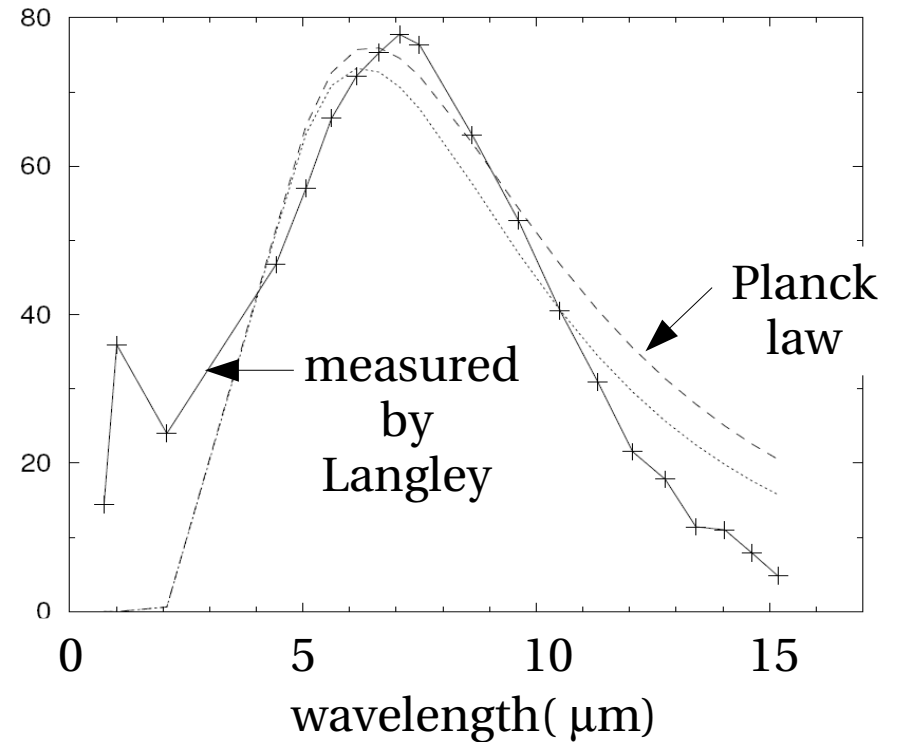
SECTION THROUGH MAIN OBSERVATORY BUILDING, 1897. LOOKING WEST.

Measurement of the infrared spectra

Measured the **solar** and **infra-red** spectrum



Measured infra-red spectra, $T=178^\circ\text{C}$



Current refractive index values are used to convert deviation angle to wavelength

- He observed, for the first time, the spectral separation between solar radiation and earth thermal radiation
- He was not satisfied by its measured of the infrared radiation coming from the moon
- (He overestimated by a factor of 2 the solar constant)

Estimate of the green house effect

On the Influence of Carbonic Acid in the Air upon the Temperature of the Ground, S. Arrhenius, 1896.

- Extended previous works done by Fourier, Pouillet, Tyndall et Langley
- **Used Langley observations to estimate the infrared absorption of the atmosphere (CO_2 , H_2O)**
- Developed the single layer greenhouse model and computed the green house effect on Earth
- CO_2 may be the cause of past climate change and may change climate in the future
- **He estimated that a doubling of $\text{CO}_2 \Rightarrow \Delta T \approx 4$ to 5°C**
- These values are close to current estimate (2-4°C)
- **Are the results obtained by Arrhenius robust** [*Ramanathan & Vogelmann, 1997*]
or fortuitous?



Svante Arrhenius

(1859-1927)

Estimate of the green house effect

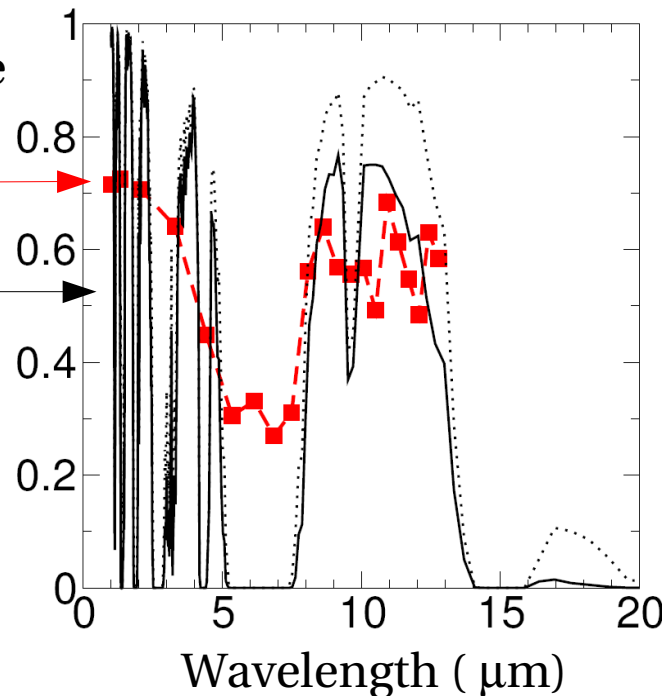
Infrared atmospheric transmittivity

$\text{CO}_2 + \text{H}_2\text{O}$

For a « standard » atmosphere

using Arrhenius data

current estimate, with a radiative model, for two water vapor contents (15 et 30 kg.m^{-2})



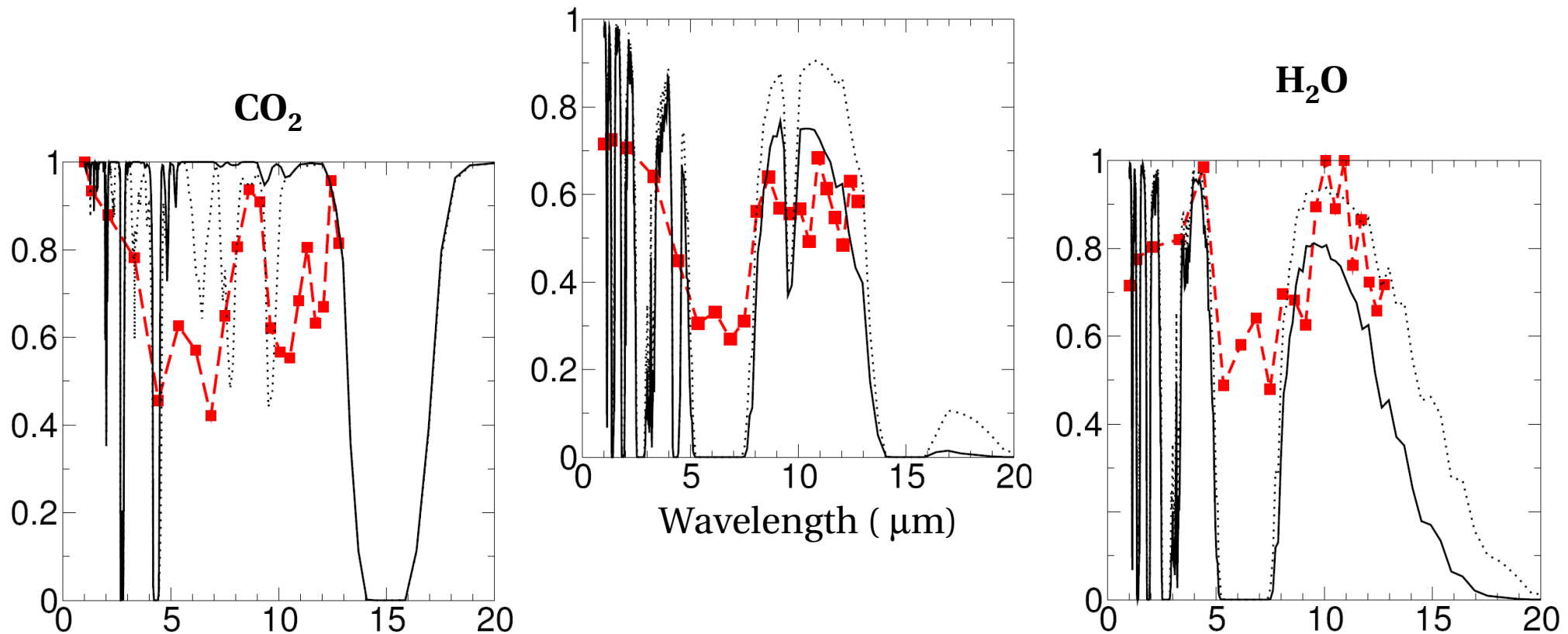
To separate the individual CO_2 and H_2O contributions, Arrhenius

- computed correlations between observed atmospheric transmittivity and humidity near surface
- made ad-hoc corrections on the observed atmospheric transmittivity

Estimate of the green house effect

Infrared atmospheric transmittivity

$\text{CO}_2 + \text{H}_2\text{O}$

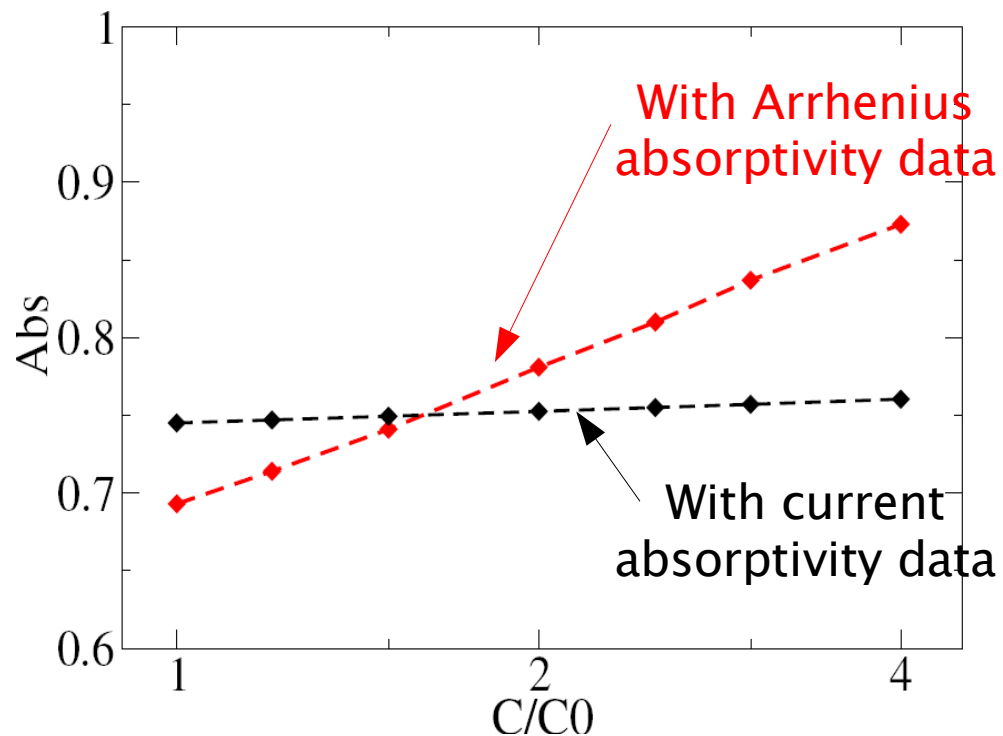


- *Observations do not include the 15 μm absorption band of CO_2*
- *The individual CO_2 and H_2O contributions are not well separate*

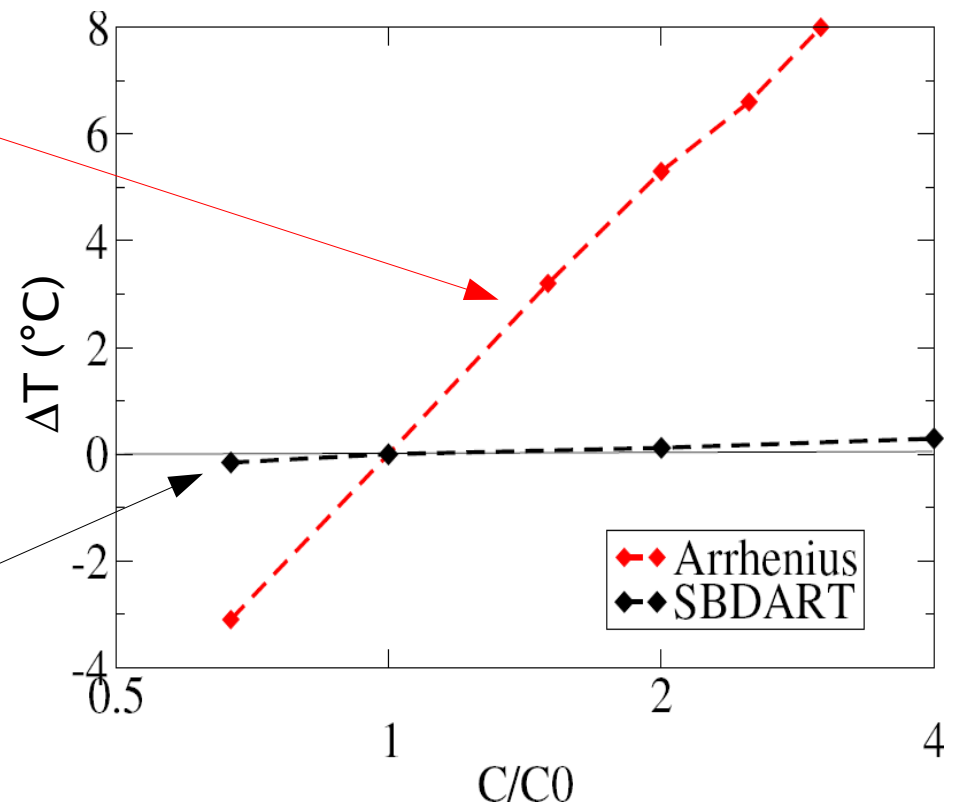
Estimate of the green house effect

Sensitivity to a change of the CO₂ concentration

Mean infrared absorptivity of the atmosphere as a function of the relative CO₂ concentration C/C_0



Temperature change ΔT (°C) as a function of the relative CO₂ concentration C/C_0 for a single layer greenhouse model



Estimate of the green house effect by Svante Arrhenius

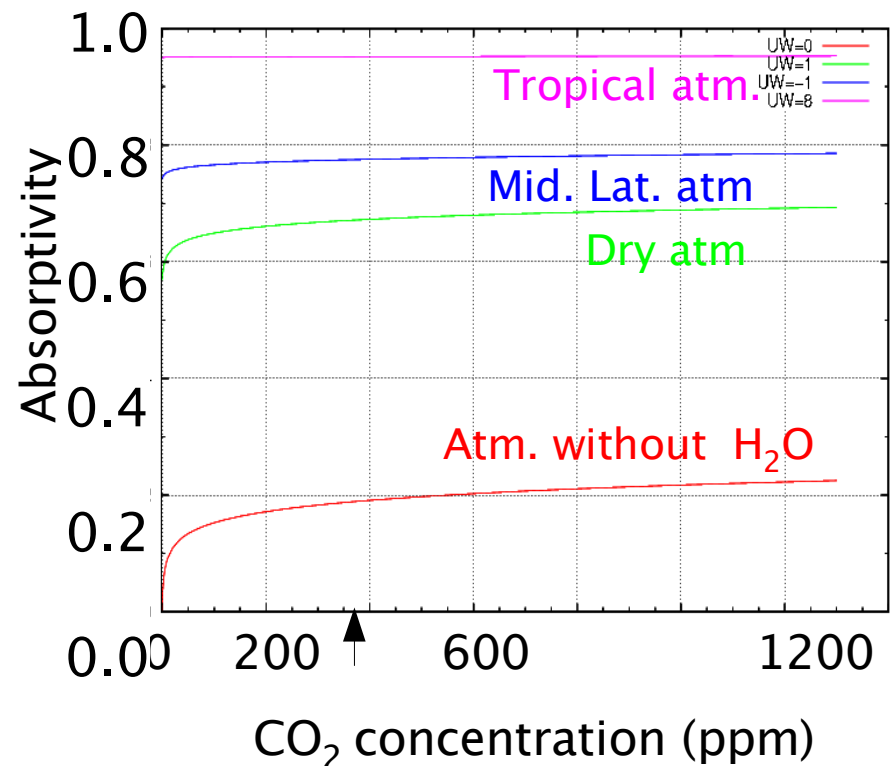
Absorptivity of the atmosphere:

- Spectral measurements do not cover the 15 μm CO_2 band
- The CO_2 absorptivity estimate is strongly contaminated by H_2O

Single layer greenhouse model:

- Do not include the vertical temperature gradient of the atmosphere
- Is not valid for greenhouse gases, like CO_2 , for which the absorption is **saturated**

The increase of temperature for a doubling of CO_2 concentration obtained by Arrhenius ($\Delta T \approx 4^\circ\text{C}$) is closed to current estimate ($\Delta T \approx 2\text{--}4^\circ\text{C}$), **but this result is fortuitous.**



Analysis of the greenhouse effect on Earth using the Net Exchange Formulation

The increase of the green house gas concentration decreases the cooling of the surface:

- But what happens within the atmosphere?
- Is the atmosphere cooled or heated?
- It is the same for H_2O and CO_2 ?

Net Exchange Formulation (NEF)

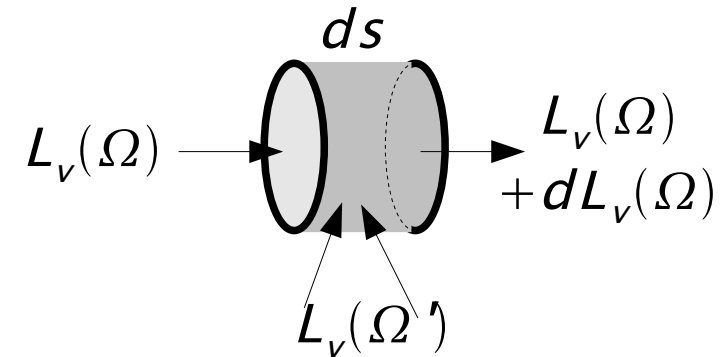
Flux formulation:

1- radiative transfer equation

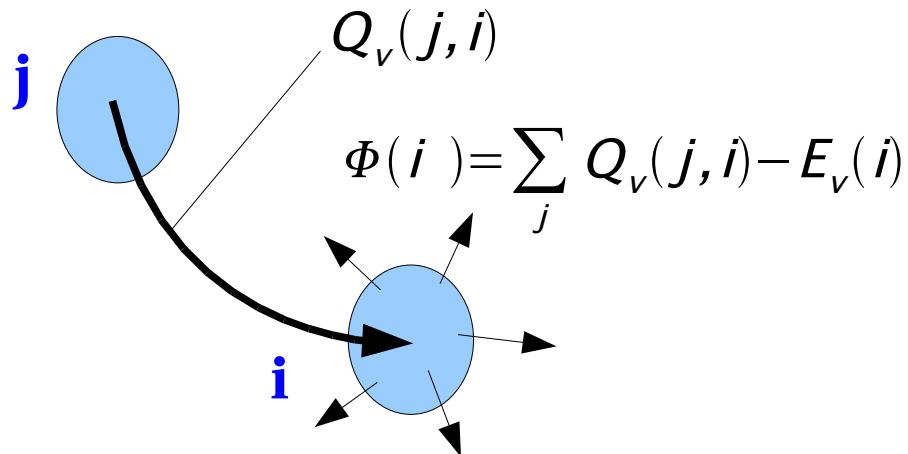
$$\frac{dL_v(\Omega)}{ds} = -\kappa_v L_v(\Omega) + \kappa_v B_v(T) - \sigma_v L_v(\Omega) + \sigma_v \frac{1}{4\pi} \int_{4\pi} P(\Omega', \Omega) L_v(\Omega') d\Omega'$$

2- computation of the radiances and radiative flux

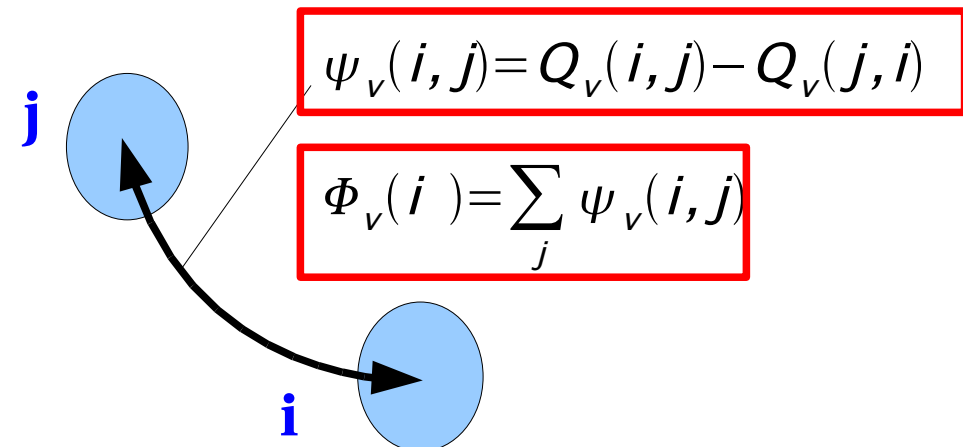
3- the heat budget Φ_v is the divergence of the radiative flux



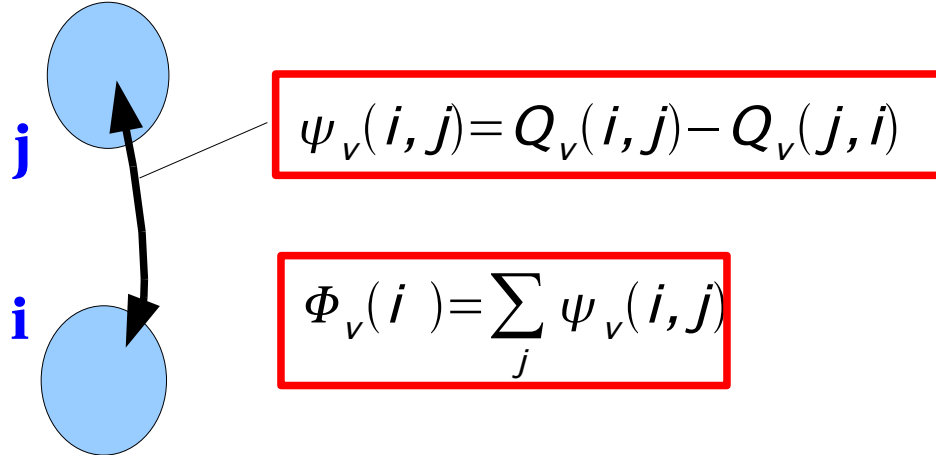
Exchange Formulation:



Net Exchange Formulation (NEF)



Net Exchange Formulation (NEF)



Net Exchange between i and j =
 difference between radiation
 emitted by i and absorbed by j
 – emitted by j and absorbed by i

$$\psi_v(i, j) = \int_{\Gamma} (\underbrace{B_v(y_y)}_{\text{Planck function}} - \underbrace{B_v(x_y)}_{\text{Optical exchange factor}}) O_v(x_y, y_y) dy$$

Planck function Optical exchange
 factor

- If the two extremities are « surfaces » (i.e. black or grey bodies, the surface or space)

$$O_v(x_y, y_y) = \tau_{v,y}$$

with τ transitivity

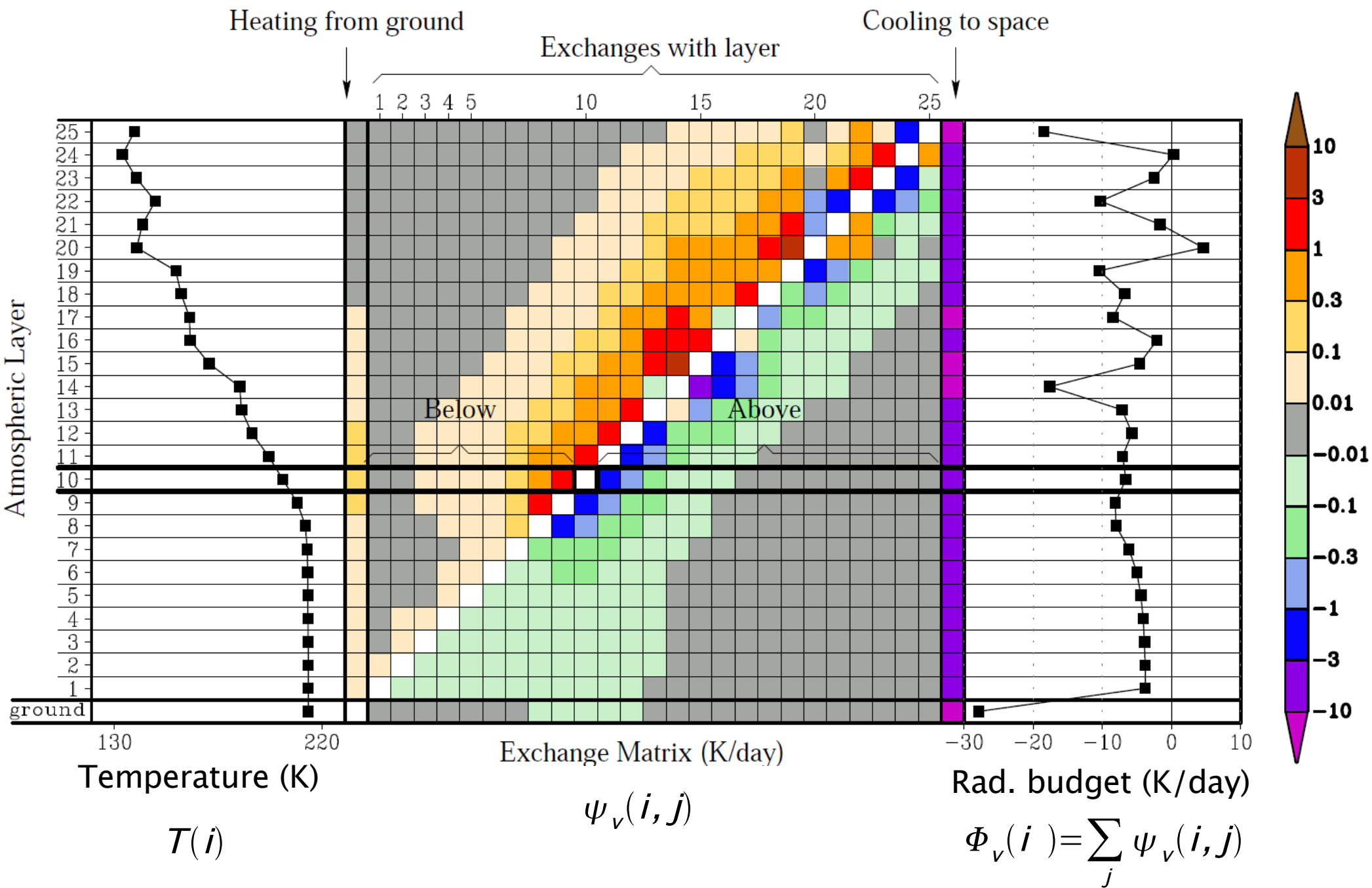
- If one extremity is a volume and the other is a « surface » (e.g. cooling to space)

$$O_v(x_y, y_y) = \left| \frac{\partial \tau_{v,y}}{\partial x_y} \right|$$

- If the two extremities are volume

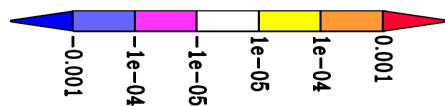
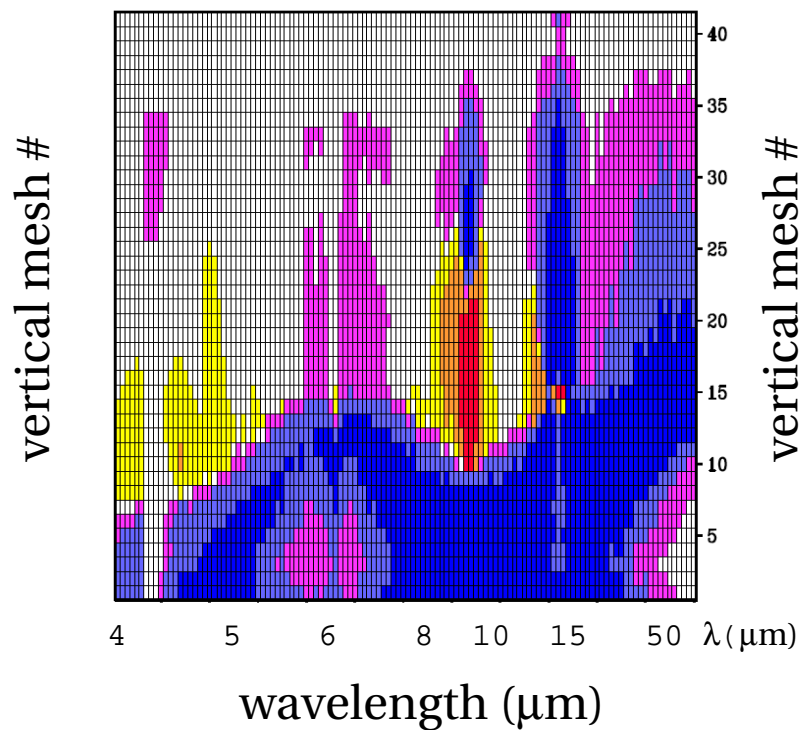
$$O_v(x_y, y_y) = \left| \frac{\partial^2 \tau_{v,y}}{\partial x_y \partial y_y} \right|$$

NEF analysis of the Mars atmosphere

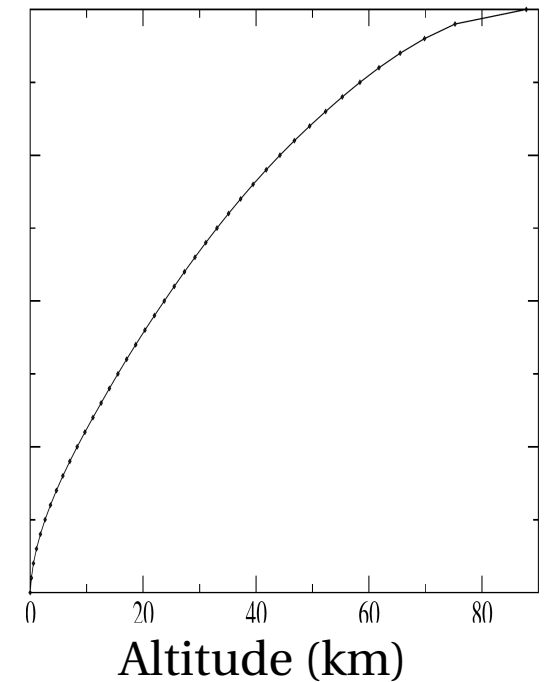
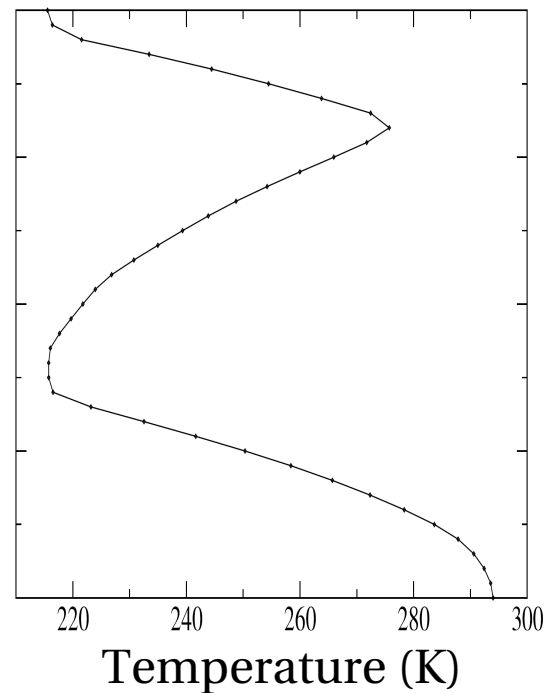


NEF analysis of the radiative exchanges for the Earth atmosphere

Spectral dependence of the **radiative budget** in the atmosphere



Mid latitude summer atmospheric profile (MLS, RTMIP case, Collins et al. 2006)



NEF analysis of the radiative exchanges

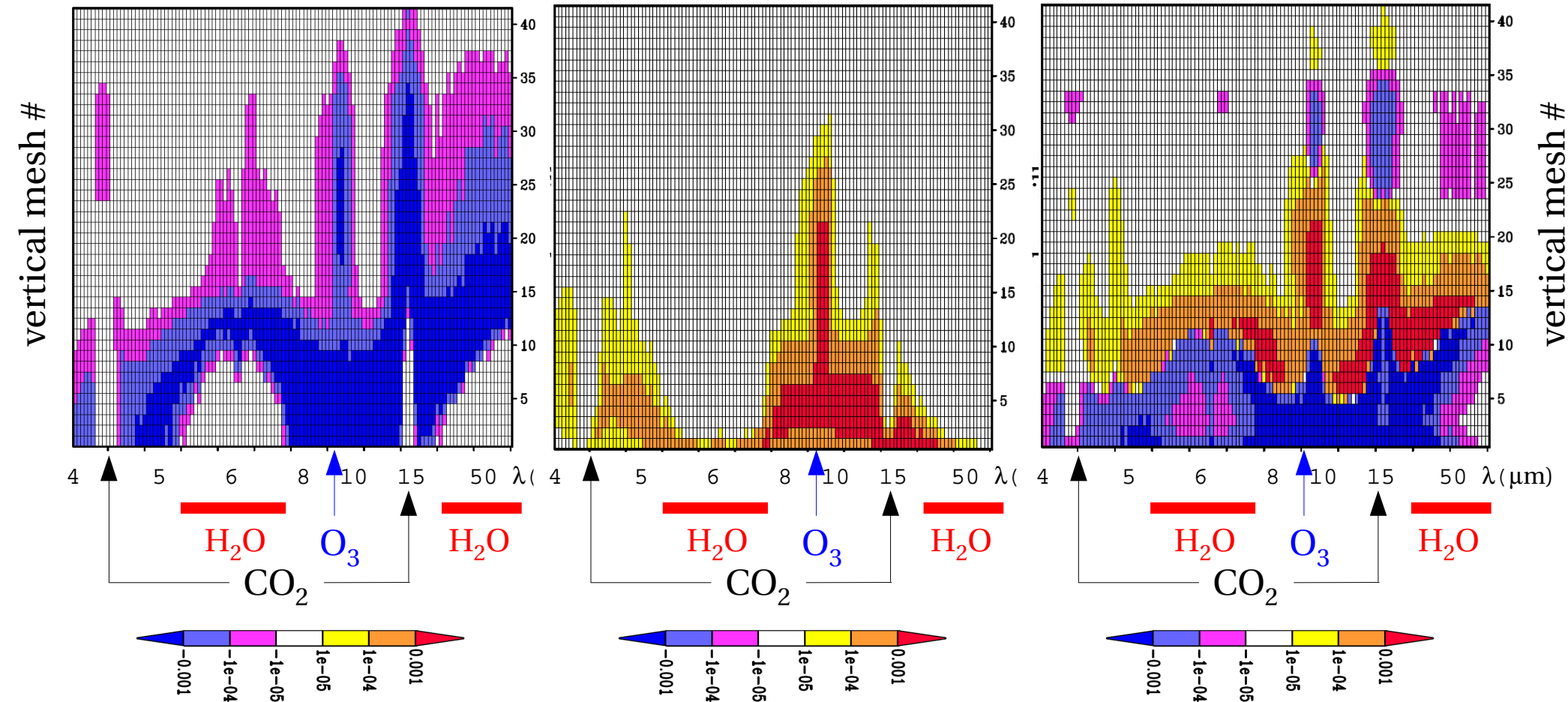
Radiative budget
of layer i

$$\Phi_i = \Psi_{i,e} + \Psi_{i,s} + \Psi_{i,a}$$

Exchange with space

Exchange with the
surface

Exchange with the rest
of the atmosphere



Focus on the exchanges with space

Net exchange between an elementary volume x of mass dm and the space s :

$$\psi_v(x, s) = -B_v(T_x) O_v(x, s) dm$$

Optical exchange factor:

$$O_v(x, s) = \left| \frac{\partial \tau_v}{\partial m} \right| = f \cdot k \cdot \exp(-f \cdot k \cdot M)$$

with:

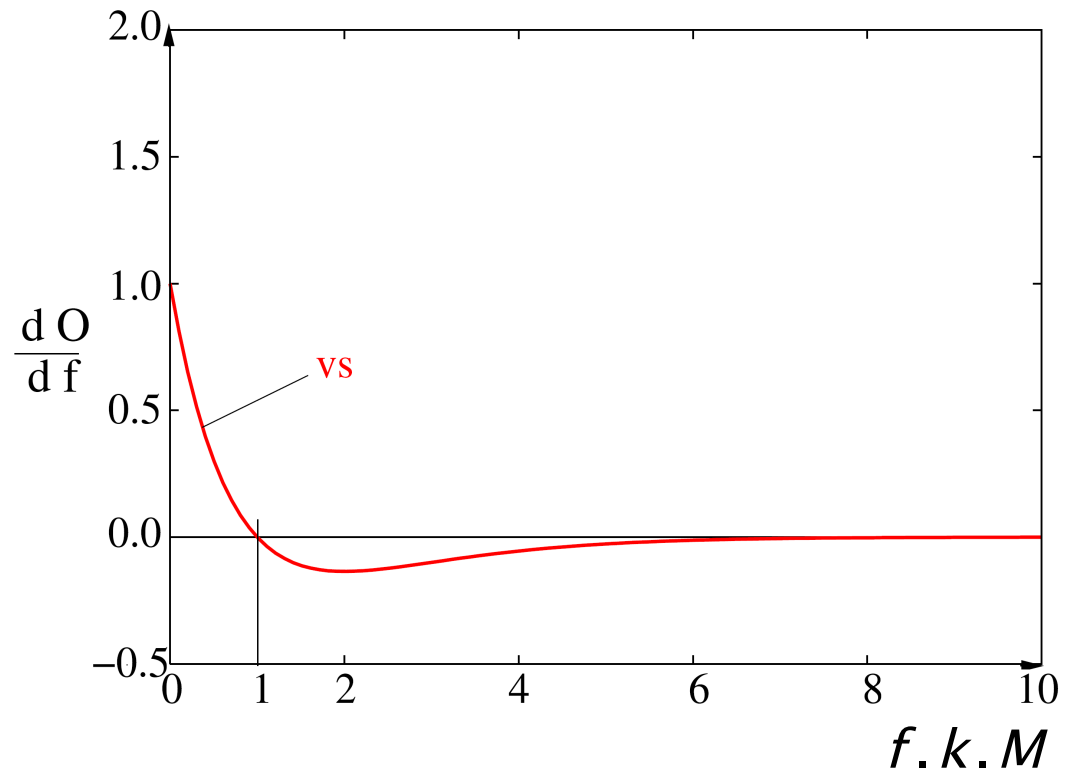
f : mass fraction of the absorbing gas

k : absorption coefficient

M : gas mass above volume x

Sensitivity of the optical exchange factor to the fraction of the absorbing gas, as a function of the optical thickness between x and the space.

The optical exchange factor increases with an increase of the fraction of the absorbing gas when the optical thickness with space is below 1, otherwise it decreases



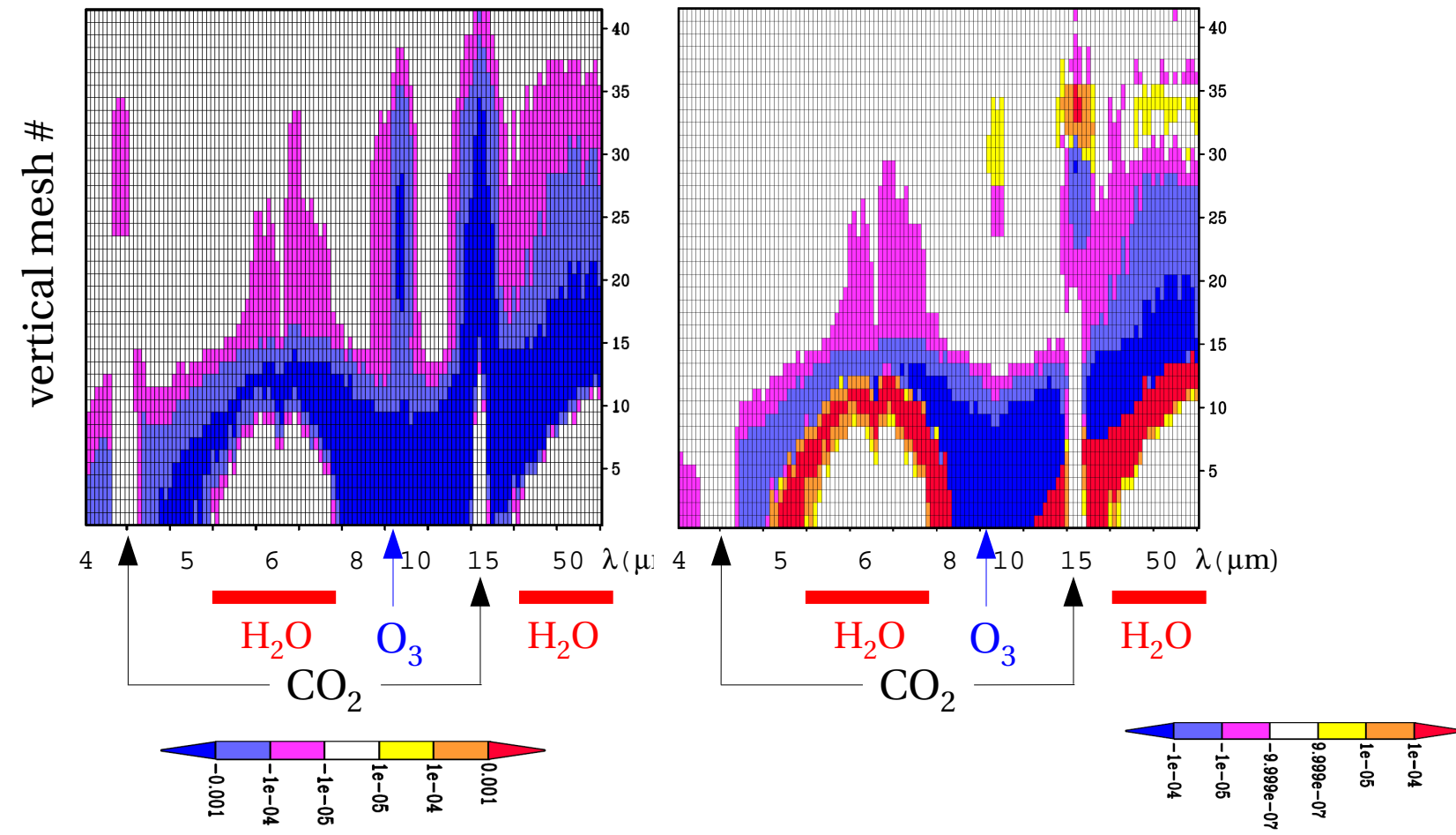
Focus on the exchanges with space

Exchanges with space

Standard atmosphere

Change of the exchanges with space

H₂O + 20%



Focus on the exchanges with space

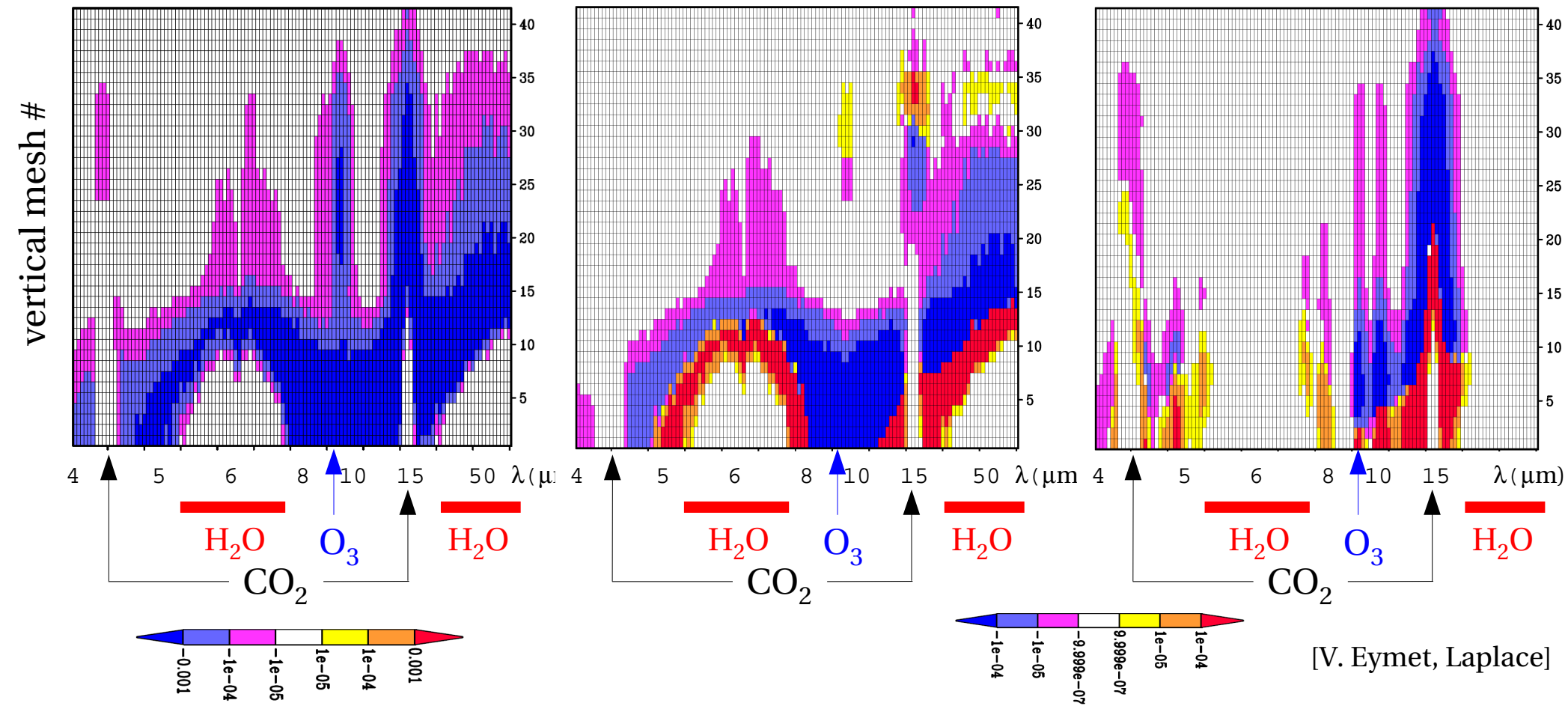
Exchanges with space

Standard atmosphere

Change of the exchanges with space

$\text{H}_2\text{O} + 20\%$

$\text{CO}_2 \times 2$



Focus on the exchanges within the atmosphere

Net exchange between two elementary volumes x and y of mass dm_x and dm_y :

$$\psi_v(x, y) = (B_v(T_y) - B_v(T_x)) O_v(x, y) dm_x dm_y$$

Optical exchange factor:

$$O_v(x, y) = \left| \frac{\partial \tau_v}{\partial m_x \partial m_y} \right| = f^2 \cdot k^2 \cdot \exp(-f \cdot k \cdot M)$$

with:

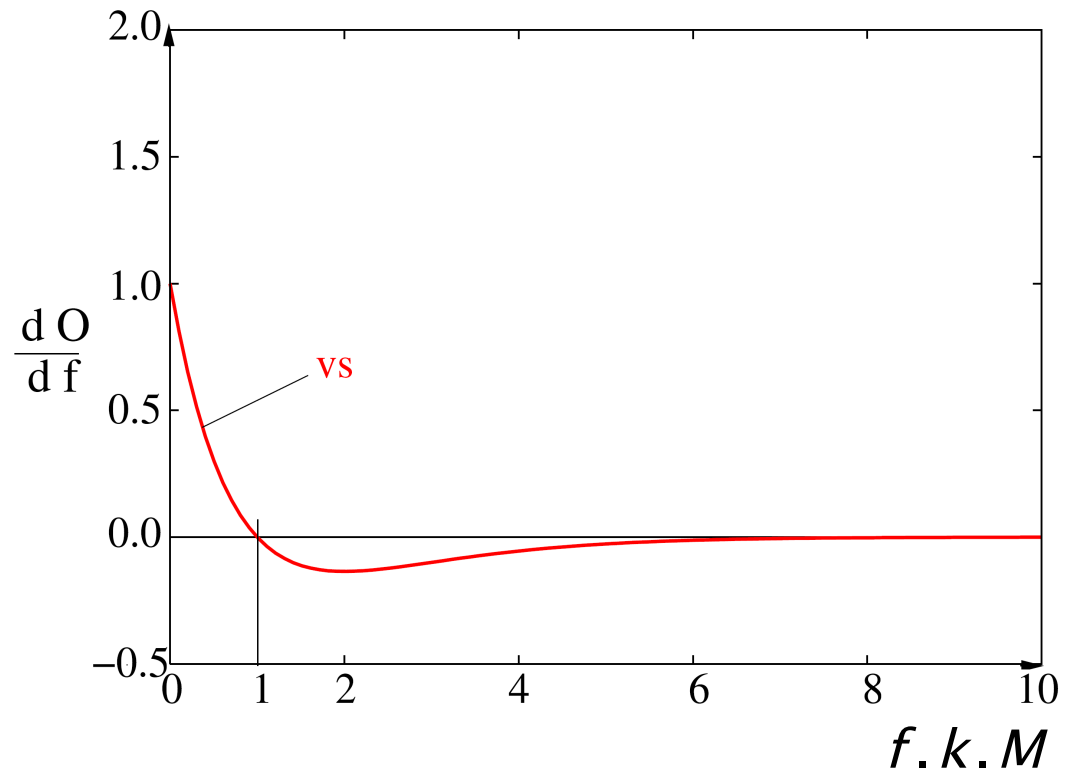
f : mass fraction of the absorbing gas

k : absorption coefficient

M : gas mass above volume x

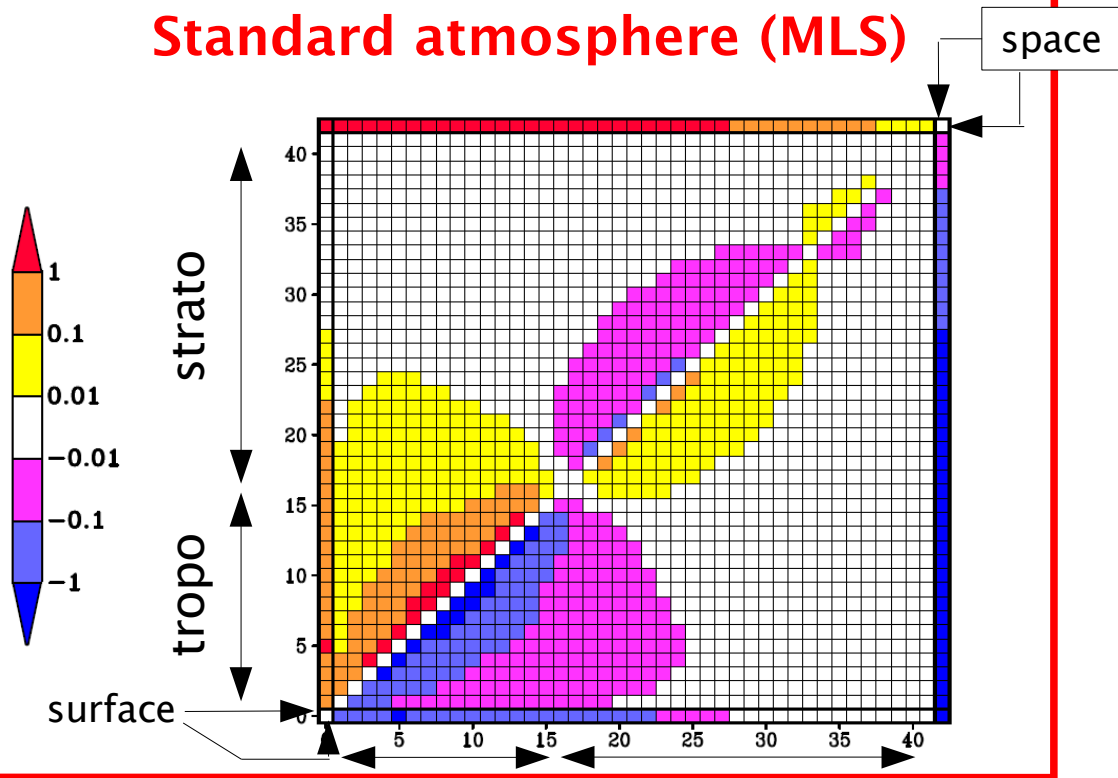
Sensitivity of the optical exchange factor to the fraction of the absorbing gas, as a function of the optical thickness between x and y

The optical exchange factor increases with an increase of the fraction of the absorbing gas when the optical thickness between x and y is below 2, otherwise it decreases



Radiative net exchange matrix ($\text{W}\cdot\text{m}^{-2}$)

Standard atmosphere (MLS)

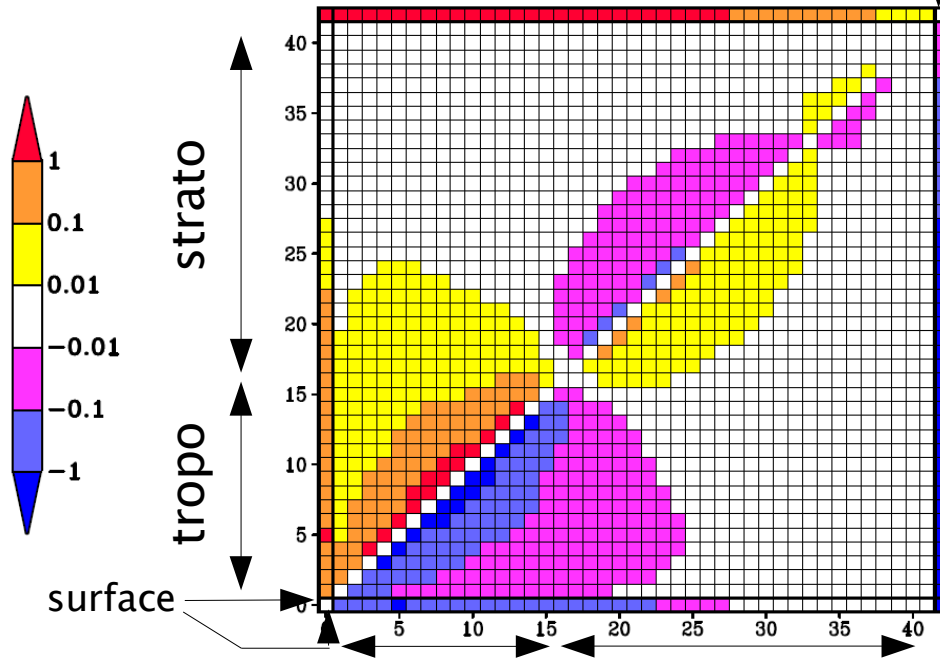


Standard atmosphere (MLS)

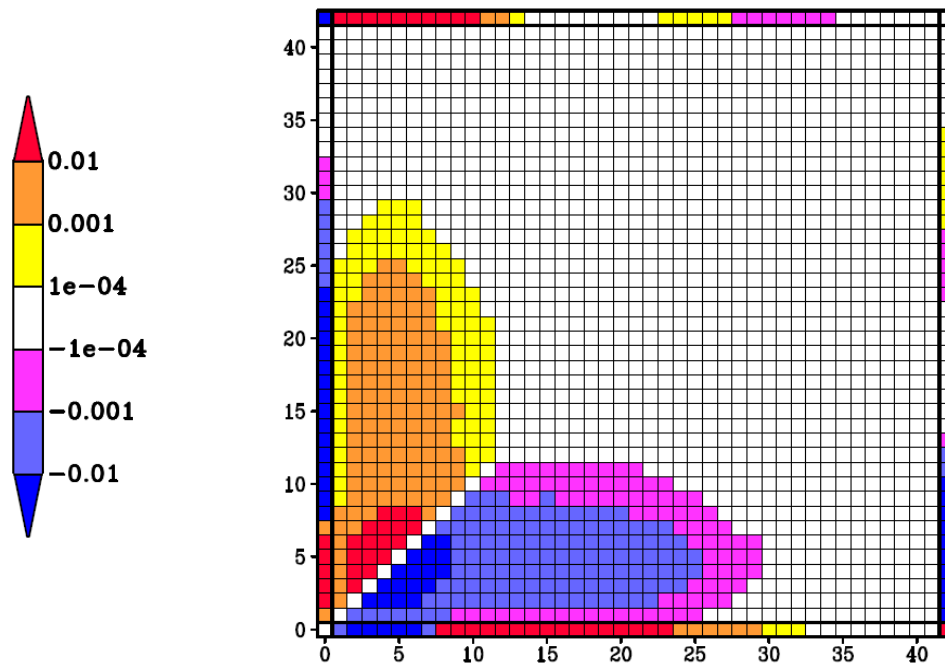
space

$\text{H}_2\text{O} + 20\%$

Radiative net exchange matrix ($\text{W}\cdot\text{m}^{-2}$)



low absorbing spectral band

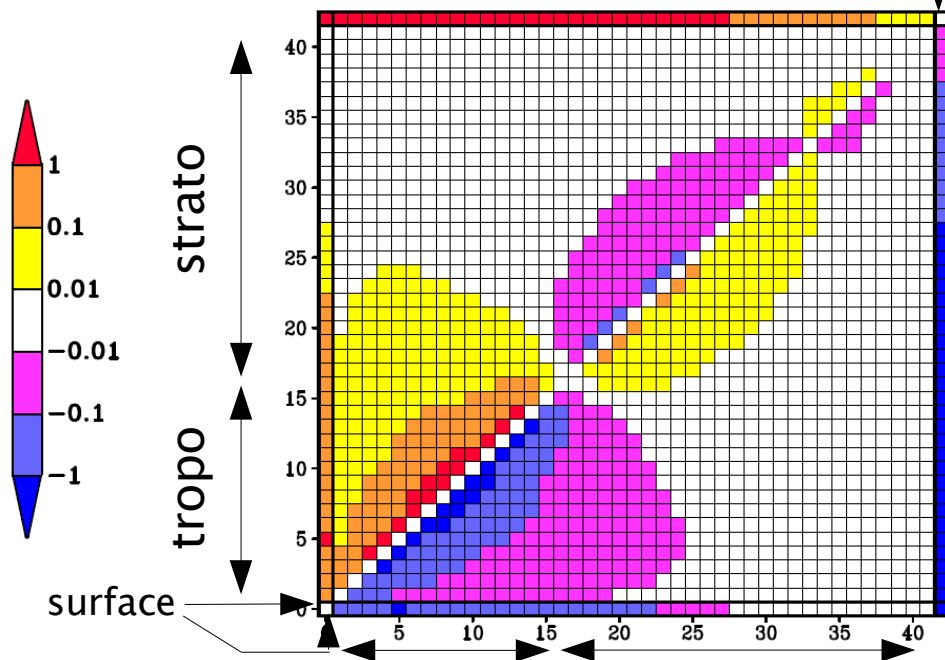


Standard atmosphere (MLS)

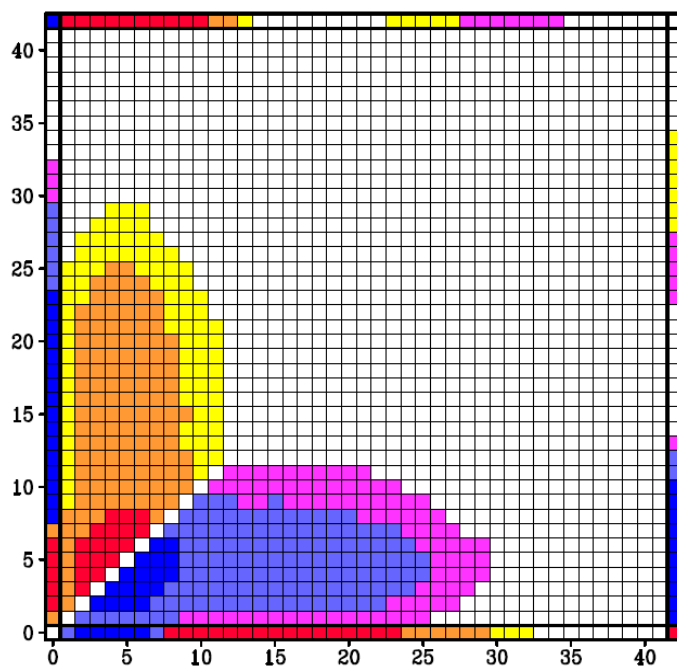
space

$H_2O + 20\%$

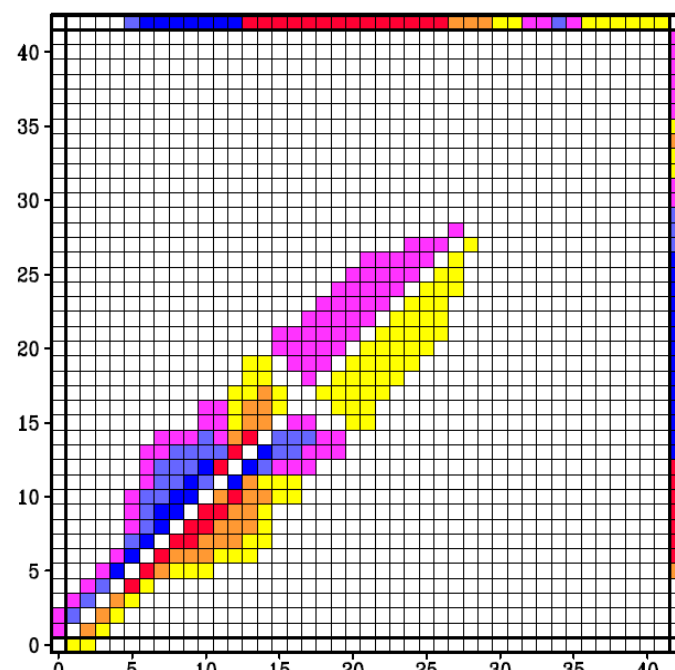
Radiative net exchange matrix ($W \cdot m^{-2}$)



low absorbing spectral band



high absorbing spectral band

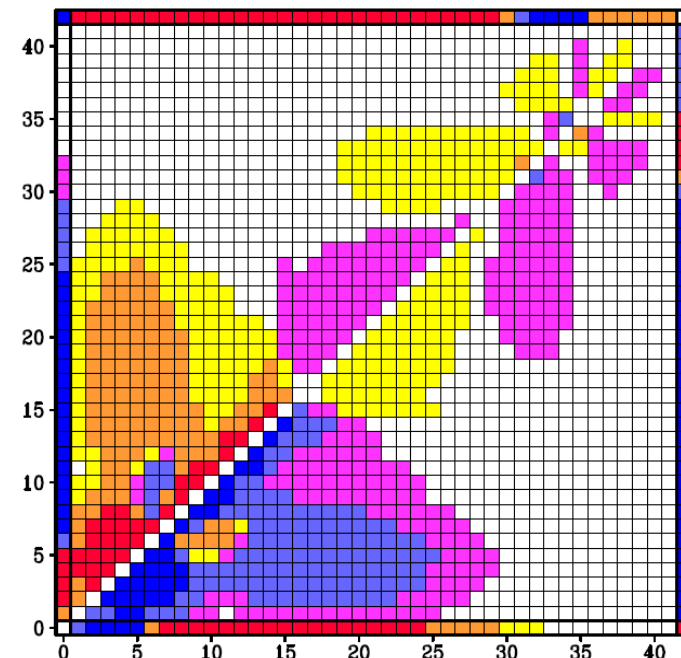
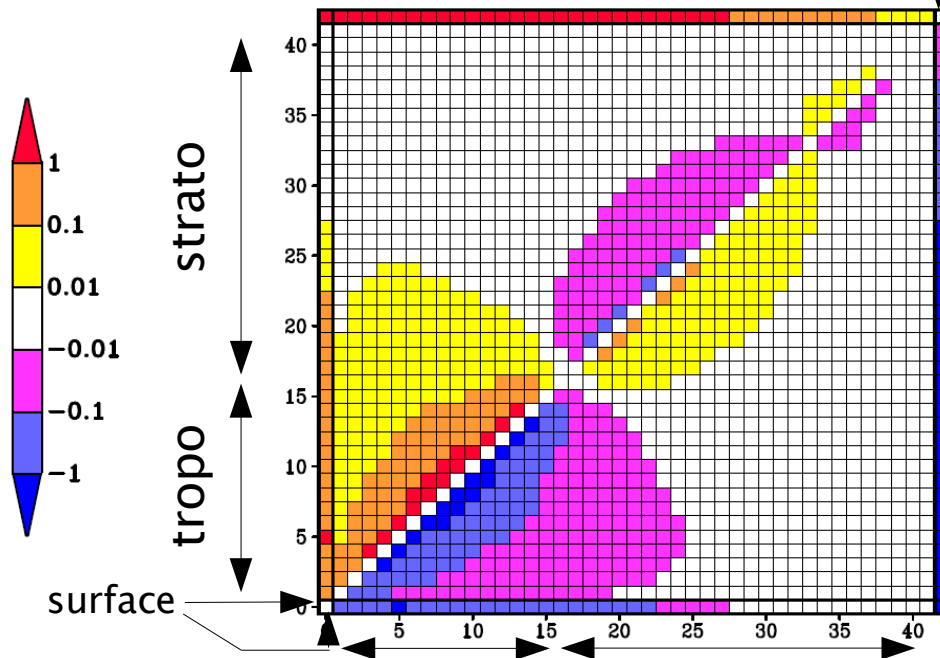


Radiative net exchange matrix ($\text{W}\cdot\text{m}^{-2}$)

Standard atmosphere (MLS)

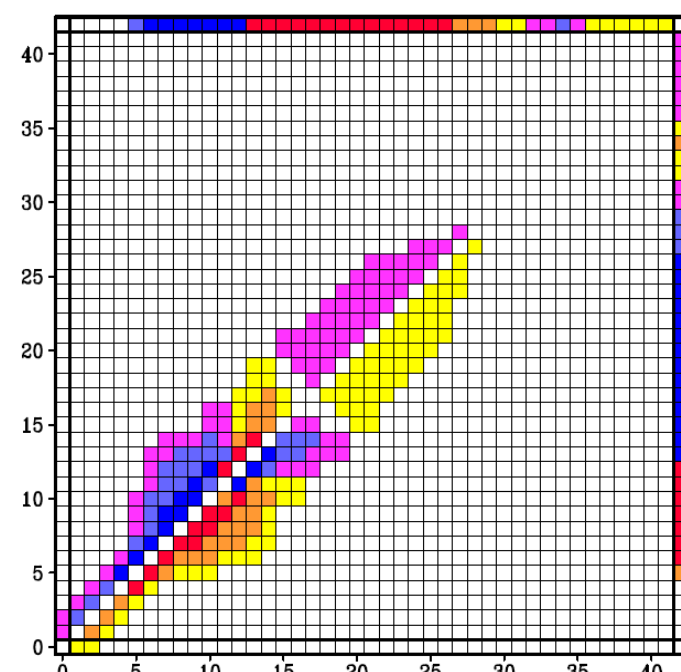
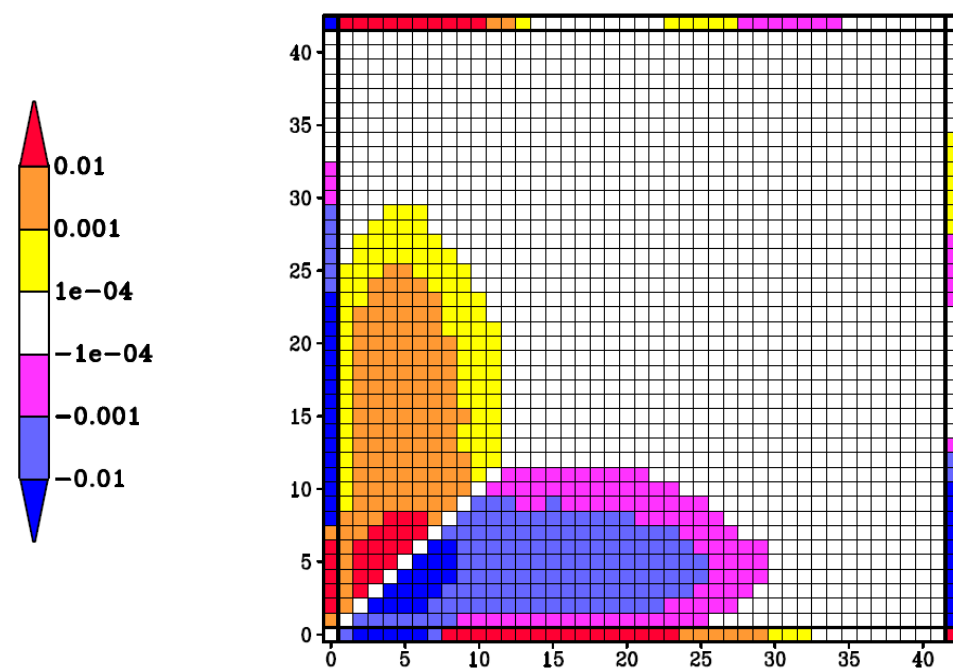
space

$\text{H}_2\text{O} + 20\%$



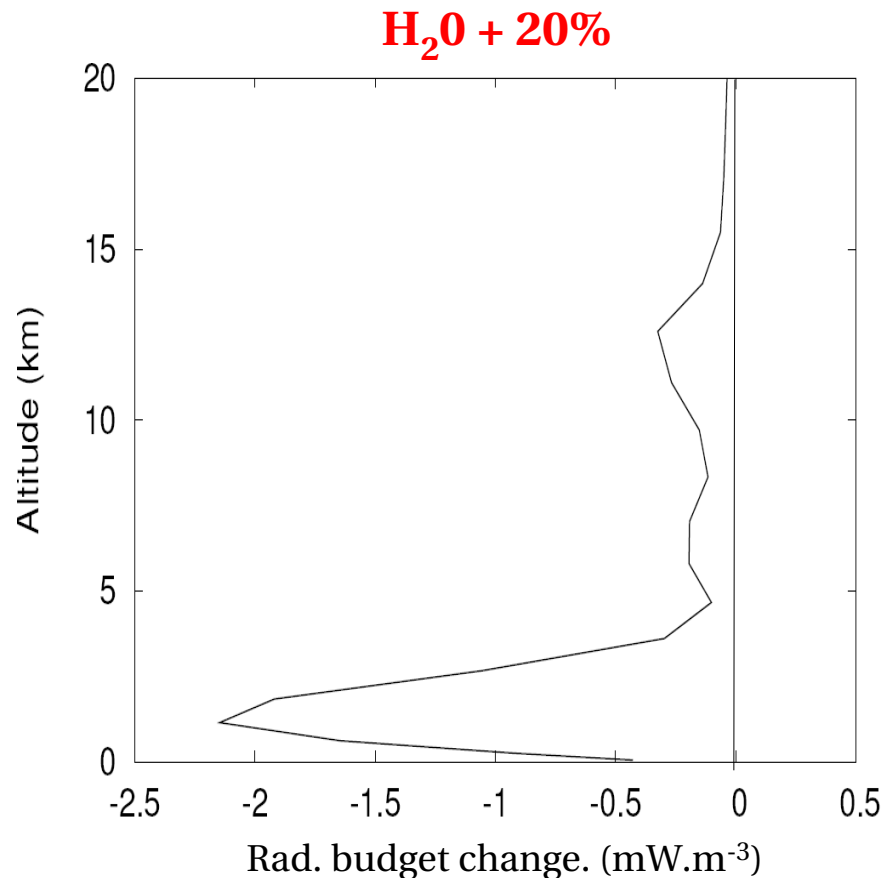
low absorbing spectral band

high absorbing spectral band



Analysis of the greenhouse effect on Earth

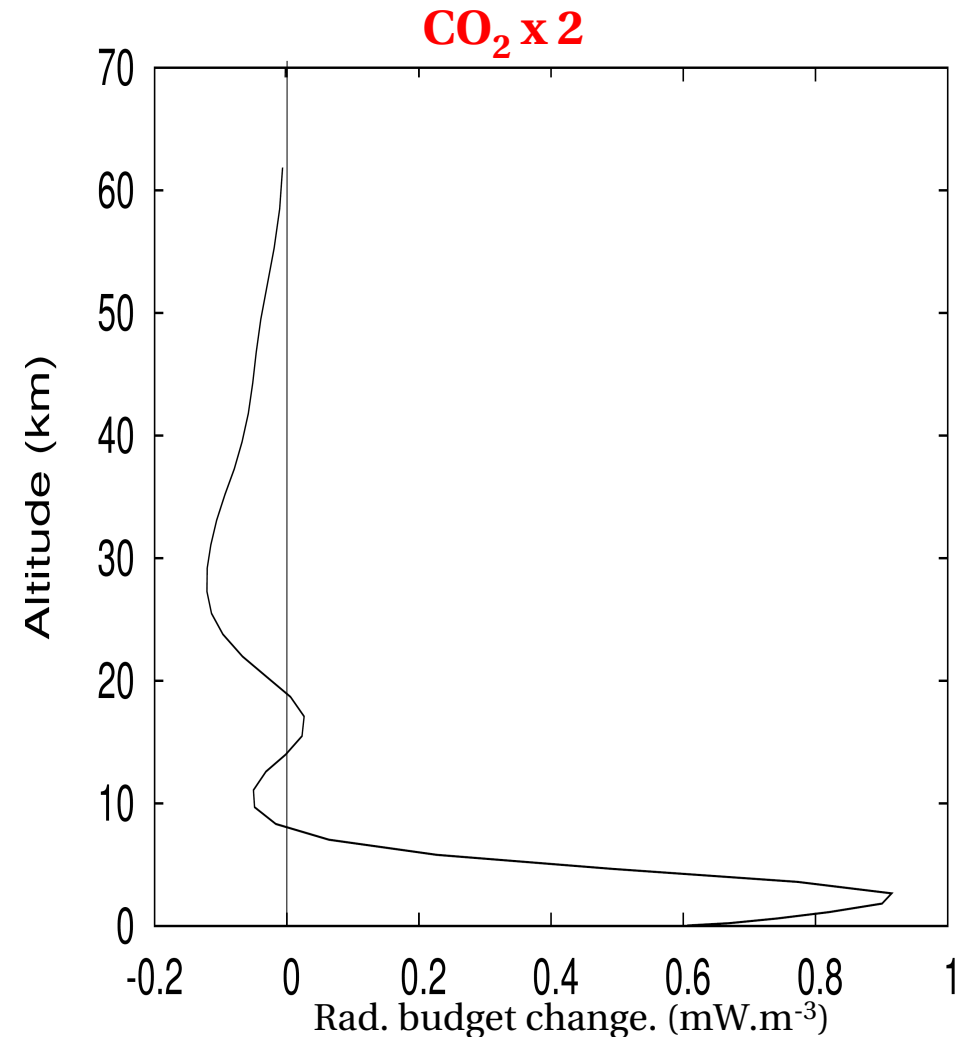
Radiative budget changes



Surface: + 12,1 W.m^{-2}

Atmosphere: - 8,1 W.m^{-2}

Space: - 4,0 W.m^{-2}



Surface: + 1,5 W.m^{-2}

Atmosphere: + 1,3 W.m^{-2}

Space: - 2,8 W.m^{-2}

In summary

Arrhenius estimate of CO₂ impact on the surface temperature is erroneous for two major reasons (that partly compensate)


- Wrong absorptivity values of CO₂ and H₂O
- Inadequacy of the single layer greenhouse model to saturated greenhouse gases, like CO₂

The consistency of Arrhenius results with current estimates appear to be fortuitous.

The Net Exchange Formalism is a powerful tool to analyse radiative exchanges.

An increase of the greenhouse gas concentration

- decreases the cooling of the surface
- cools or heats the atmosphere, depending on the radiative properties of the gas. On average:
 - for H₂O: increase of the cooling of the atmosphere
 - for CO₂: decrease of the cooling of the atmosphere

A vibrant rainbow arches across a cloudy sky, with a line of trees visible at the bottom. The text "Thank you for your attention" is centered in the middle of the image.

Thank you for your attention